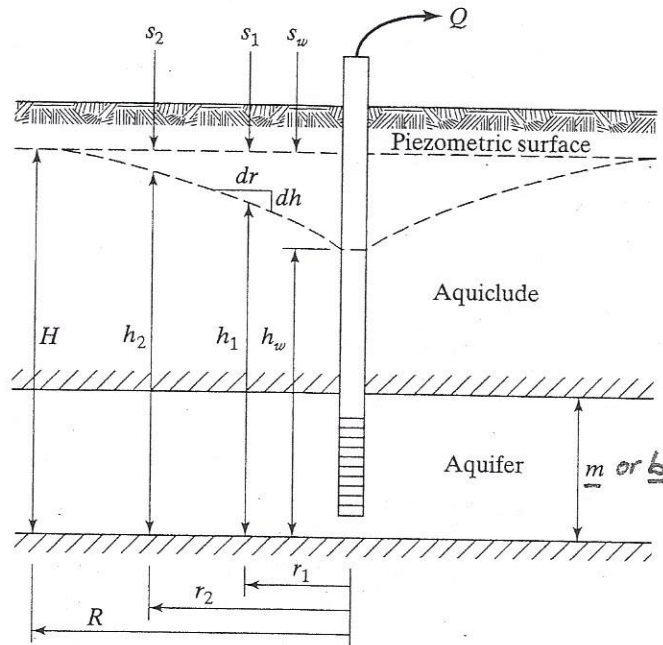


## Steady-State Groundwater Flow:

### Single Well Pumping in Confined Aquifer – Thiem's Equation



$$Q = \frac{2\pi T(h_2 - h_1)}{\ln(r_2/r_1)}$$

$$s_w = H - h_w = \frac{Q}{2\pi T} \ln(R/r_w)$$

#### Example 9.7

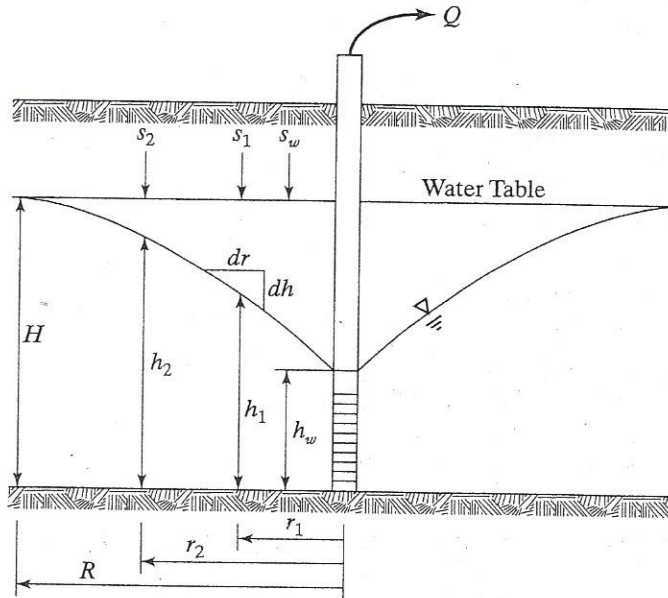
##### Steady-State Drawdown in a Confined Aquifer

A 24-inch diameter well is installed in a confined aquifer using a 6-inch gravel pack ( $r_w = 1.5$  ft). The aquifer has a transmissivity of 40,000 gpd/ft, and the well is pumped at a rate of 2,000 gpm. If the radius of influence ( $R$ ) of the well is 100,000 ft, determine the steady-state drawdown at the well.

$$s_w = \frac{Q}{2\pi T} \ln\left(\frac{R}{r_w}\right) = \frac{2,000 \times 1,440}{2\pi \times 40,000} \ln\left(\frac{100,000}{1.5}\right) = 127 \text{ ft}$$

## Steady-State Groundwater Flow:

### Single Well Pumping in Unconfined Aquifer – Dupuit's Equation



$$Q = \frac{\pi K(h_2^2 - h_1^2)}{\ln(r_2/r_1)}$$

#### Example 9.8

##### Steady-State Drawdown in an Unconfined Aquifer

A 12-inch diameter well is installed in an unconfined aquifer with a saturated thickness of 100 ft. The aquifer has a transmissivity of 40,000 gpd/ft, and the well is pumped at a rate of 600 gpm. If the radius of influence ( $R$ ) of the well is 4,000 ft, determine the steady-state drawdown at the well.

$$h_2^2 - h_1^2 = \frac{Q}{\pi K} \ln\left(\frac{r_2}{r_1}\right)$$

$$100^2 - h_w^2 = \frac{600 \times 1,440}{3.1416 \times 400} \ln\left(\frac{4,000}{0.5}\right)$$

$$h_w^2 = 10,000 - 6,179$$

$$h_2 = 61.8 \text{ ft}$$

$$s_w = h_2 - h_w = 38.2 \text{ ft}$$

For comparison, the computed drawdown using the confined aquifer equation is

$$s_w = \frac{Q}{2\pi T} \ln\left(\frac{R}{r_w}\right) = \frac{600 \times 1,440}{2\pi \times 40,000} \ln\left(\frac{4,000}{0.5}\right) = 31.0 \text{ ft}$$